Calculating Surface Water Time of Travel to Protect Public Water Intakes in West Virginia

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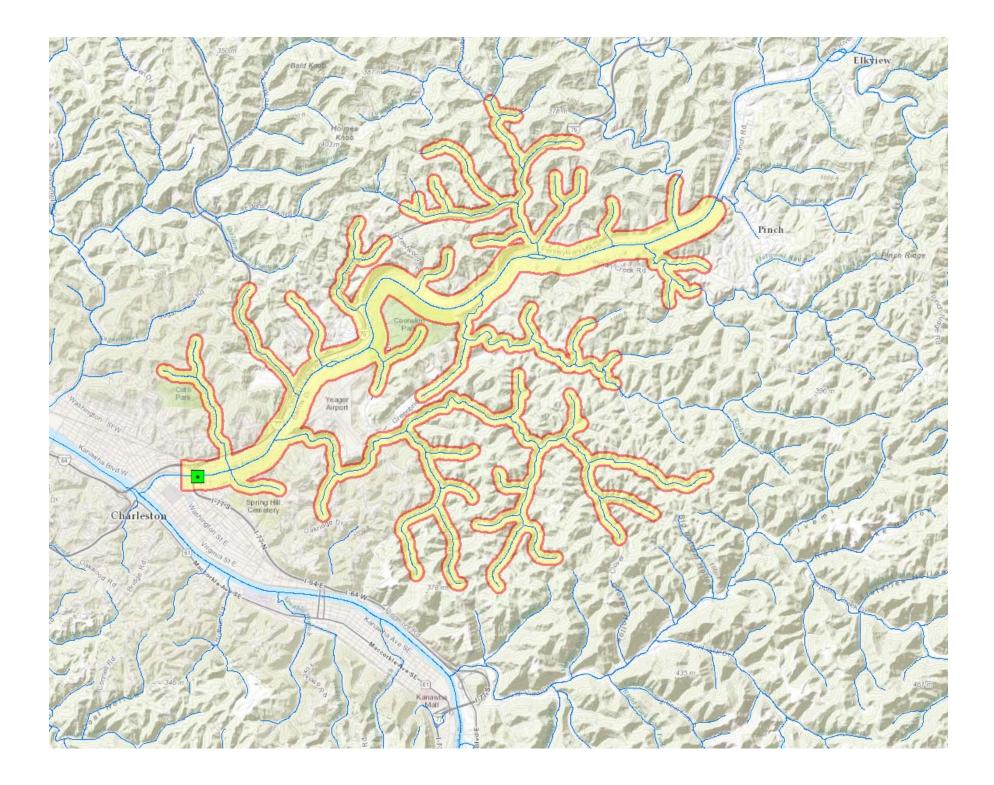
School of Natural Resources

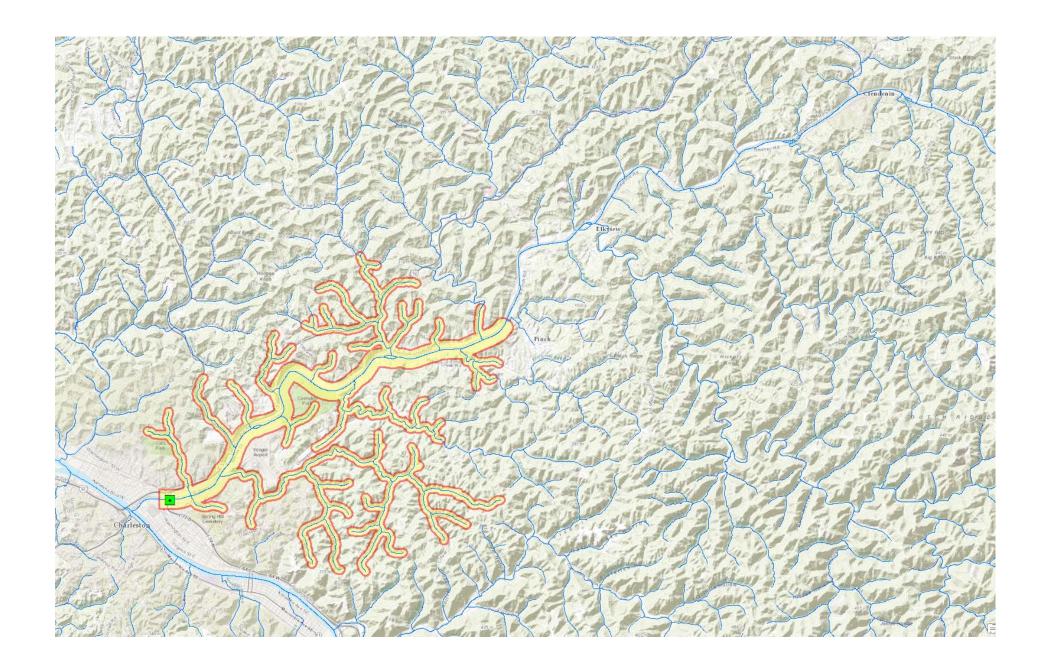


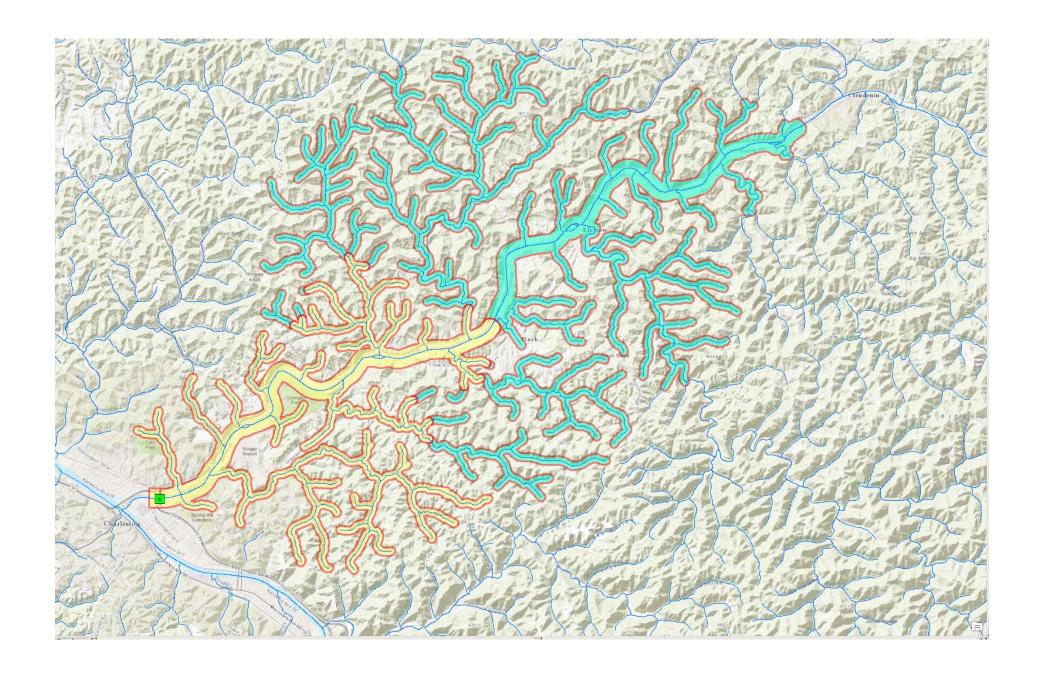
Zones of Critical Concern

- 5 hour travel time above source water intakes
- 10 hour for Zones of Peripheral Concern
- Mainstem buffer of 1,000 ft
- Tributary buffer of 500 ft

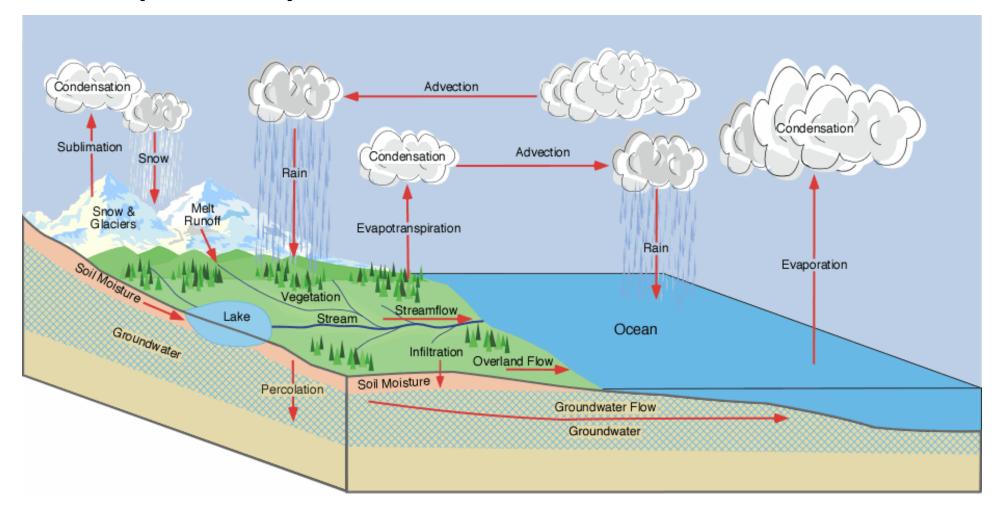
165 completed sites as of Oct 1, 2015 with ZCC and ZPCs which include Ohio River Intakes which follow different delineation rules





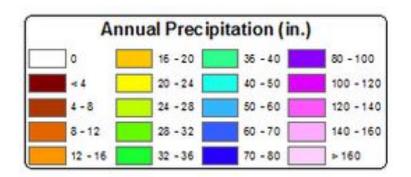


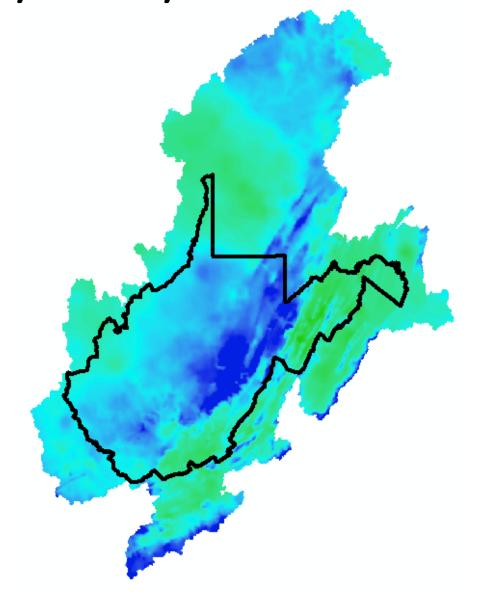
Hydro Cycle



Controls on the Hydro Cycle

Climate





Controls on the Hydro Cycle

∐G [m]

• Morphology

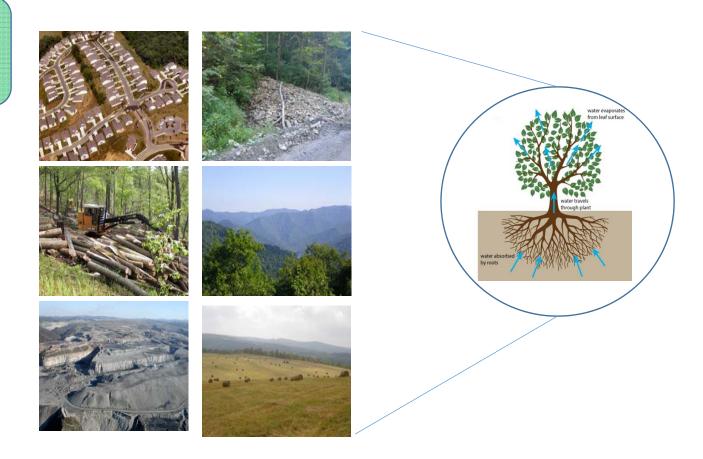
relict terrace

N

100 200 0 400 Meters

Controls on the Hydro Cycle

Landcover



Prediction of Traveltime and Longitudinal Dispersion in Rivers and Streams

By Harvey E. Jobson

U.S. GEOLOGICAL SURVEY
Water-Resources Investigations Report 96-4013



Reston, Virginia 1996

Prediction of Traveltime and Longitudinal Dispersion in Rivers and Streams

by Harvey E. Jobson

Abstract

The possibility of a contaminant being accidentally or intentionally spilled upstream from a water supply is a constant concern to those diverting and using water from streams and rivers. Although many excellent models are available to estimate traveltime and dispersion, none can be used with confidence before calibration and verification to the particular river reach in question. Therefore, the availability of reliable input information is usually the weakest link in the chain of events needed to predict the rate of movement, dilution, and mixing of contaminants in rivers and streams.

Measured tracer-response curves produced from the injection of a known quantity of soluble tracer provide an efficient method of obtaining the necessary data. The purpose of this report is to use previously presented concepts along with extensive data collected on time of travel and dispersion to provide guidance to water-resources managers and planners in responding to spills. This is done by providing methods to estimate (1) the rate of movement of a contaminant through a river reach, (2) the rate of attenuation of the peak concentration of a conservative contaminant with time, and (3) the length of time required for the contaminant plume to pass a point in the river. Although the accuracy of the predictions can be greatly increased by performing time-of-travel studies on the river reach in question, the emphasis of this report is on providing methods for making estimates where few data are available.

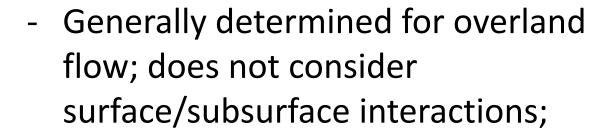
Results from rivers of all sizes can be combined by defining the unit concentration as that concentration of a conservative pollutant that would result from injecting a unit of mass into a unit of flow. Unit-peak concentrations are compiled for more than 60 different rivers representing a wide range of sizes, slopes, and geomorphic types. Analyses of these data indicate that the unit-peak concentration is well correlated with the time required for a pollutant cloud to reach a specific point in the river. The variance among different rivers is, of course, larger than for a specific river reach. Other river characteristics that were compiled and included in the correlation included the drainage area, the reach slope, the mean annual discharge, and the discharge at the time of the measurement. The most significant other variable in the correlation was the ratio of the river discharge to mean annual discharge.

The prediction of the traveltime is more difficult than the prediction of unit-peak concentration; but the logarithm of stream velocity can be assumed to be linearly correlated with the logarithm of discharge. More than 980 subreaches for about 90 different rivers were analyzed and prediction equations were developed based on the drainage area, the reach slope, the mean annual discharge, and the discharge at the time of the measurement. The highest probable velocity, which will result in the highest concentration, is usually of concern after an accidental spill. Therefore, an envelope curve for which more than 99 percent of the velocities were smaller was developed to address this concern.

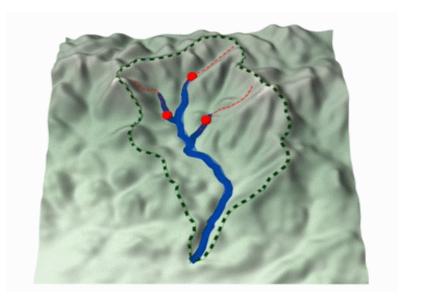
The time of arrival of the leading edge of the pollutant indicates when a problem will first exist and defines the overall shape of the tracer-response function. The traveltime of the leading edge is generally about 89 percent of the traveltime to the peak concentration.

<u>Travel time</u> – time it takes a unit volume of water to travel between two points;



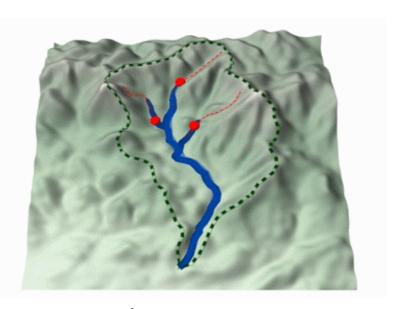


 Therefore, sheetflow moving in response to gradient (topography).



Jobson, 1996 determined that unit-peak concentration of tracer, C_{uo} :

$$C_{up} = 857T_p^{-0.760(Q/Q_a)^{-0.079}}$$



 T_p = elapsed time from injection to peak concentration, hrs.

Q= event streamflow, ft³/s

 Q_a = mean annual streamflow, ft³/s

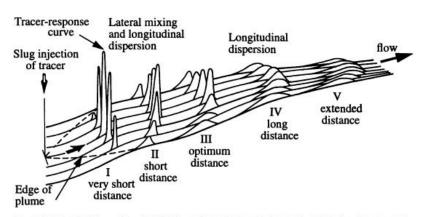


Figure 1. Lateral mixing and longitudinal dispersion patterns and changes in distribution of concentration downstream from a single, center, slug injection of tracer. (Modified from Kilpatrick, 1993, p. 2.)

Jobson ,1996 – determined velocity of peak concentration could be estimated as f(drainage area, streamflow, & slope)

Developed for 980 sub-reaches based on drainage area, reach slope, mean annual streamflow, and streamflow at time of measurement;

Log of stream velocity is assumed to be linearly correlated with log (Q);

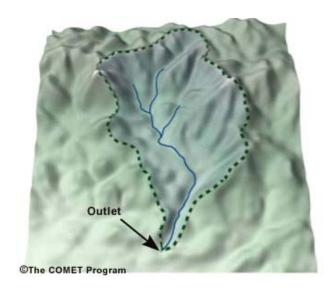
Developed for free-flowing streams, not applicable to regulated streams.

$$V_p = 0.308 + \left(0.0143 * D_a^{\prime 0.919} Q_a^{\prime -0.469} S^{0.159} \frac{Q}{D_a}\right)$$

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$$D_a'$$
 Drainage area index:

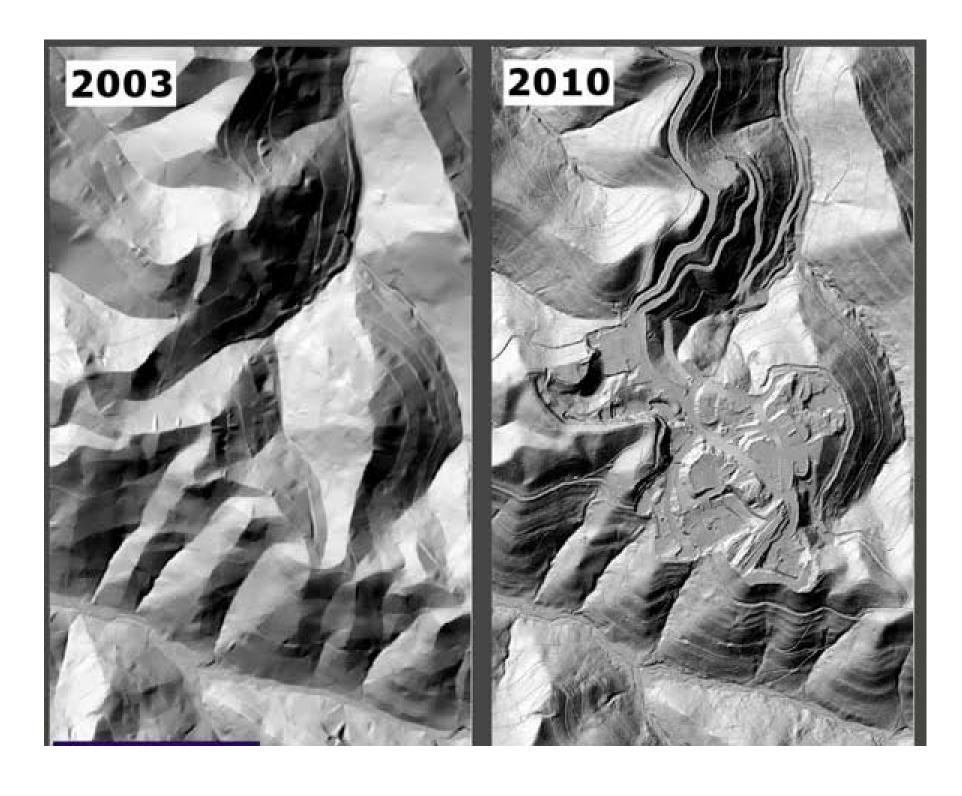
$$D_a'$$
 Drainage area index: $D_a' = \frac{D_a^{1.25} \ g^{0.5}}{Q_a}$



 D_a : drainage area; DEM-driven; updated using the 2010 WV Lidar data DEM & 2003 SAMB DEM;

g: acceleration due to gravity – constant;

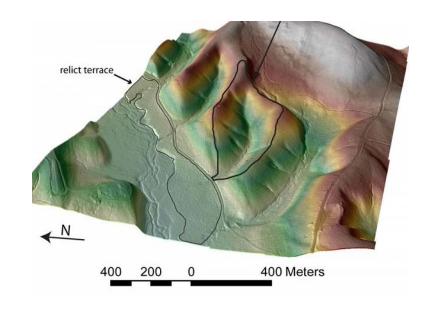
 Q_a : mean annual streamflow;

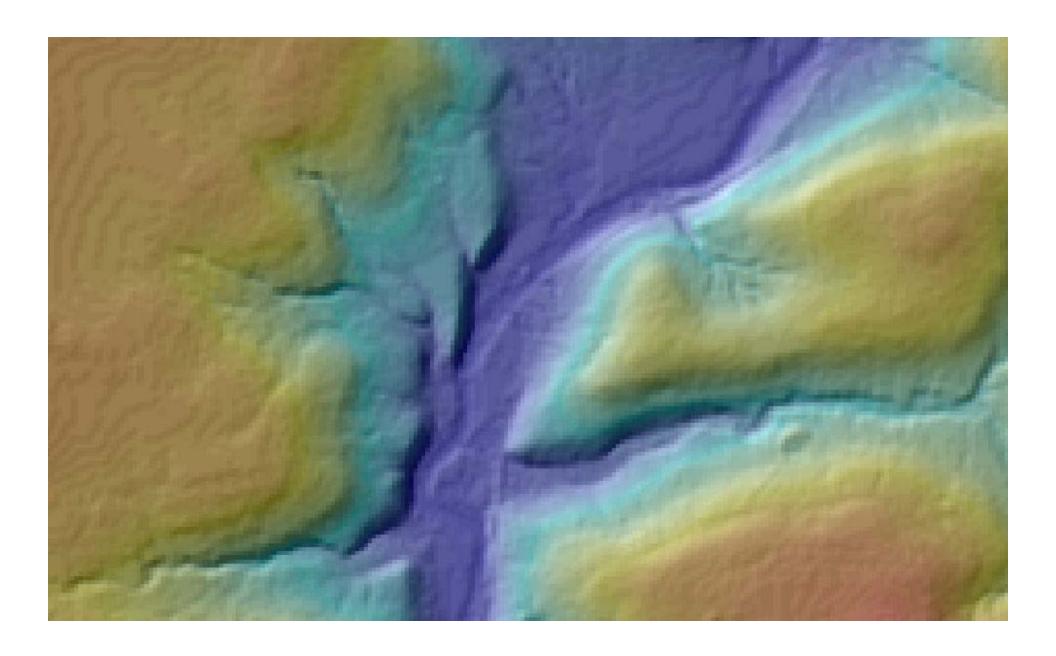


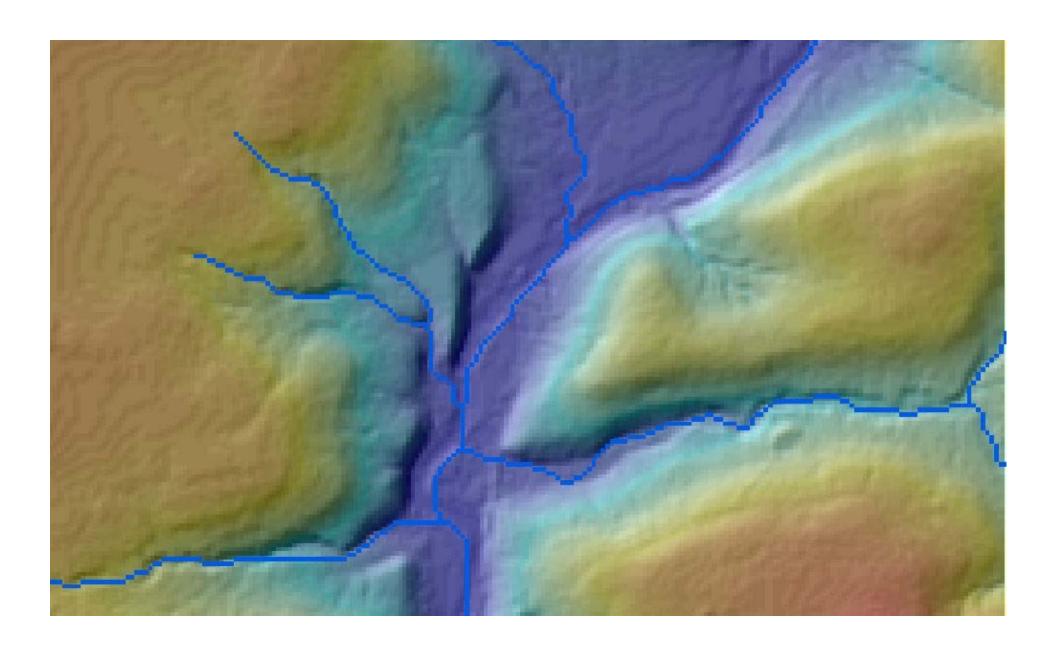
$$V_p = 0.308 + \left(0.0143 * D_a^{\prime 0.919} \ Q_a^{\prime -0.469} \ S^{0.159} \ \frac{Q}{D_a}\right)$$

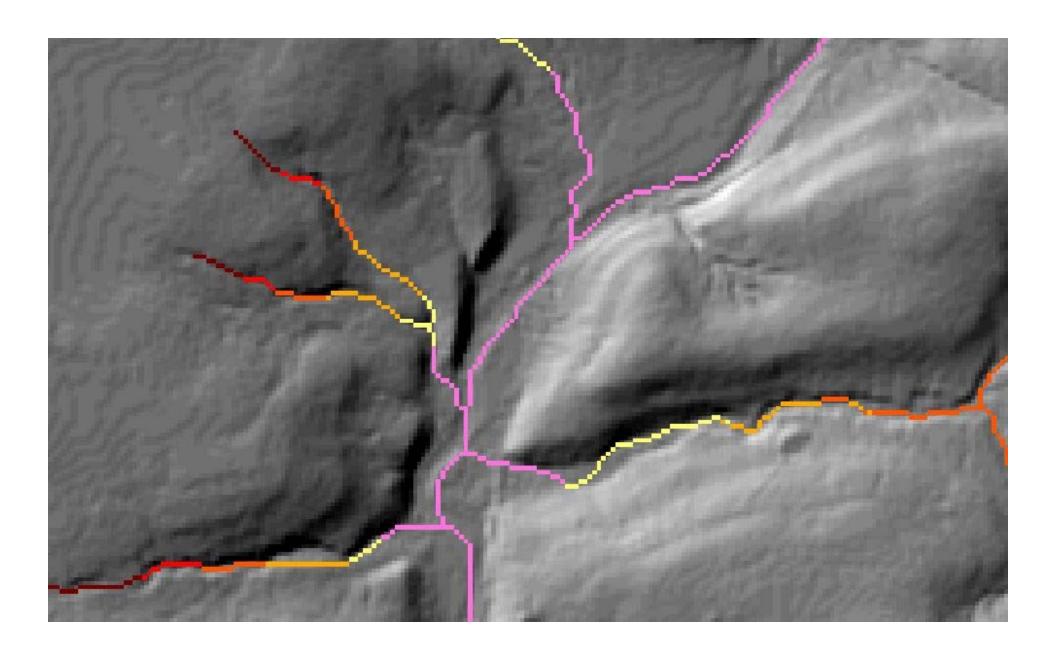
S, Slope:

Calculated similar to previous version, but uses updated DEM





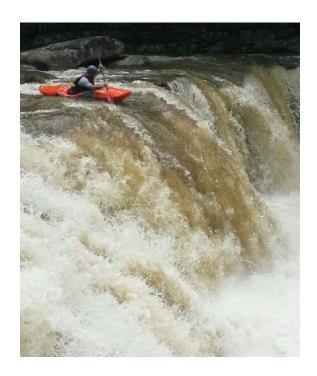




$$V_p = 0.308 + \left(0.0143 * D_a^{\prime 0.919} Q_a^{\prime -0.469}\right) S^{0.159} \frac{Q}{D_a}$$

 Q_a' Dimensionless streamflow index:

$$Q_a' = \frac{Q}{Q_a}$$



 Q_q : mean annual streamflow

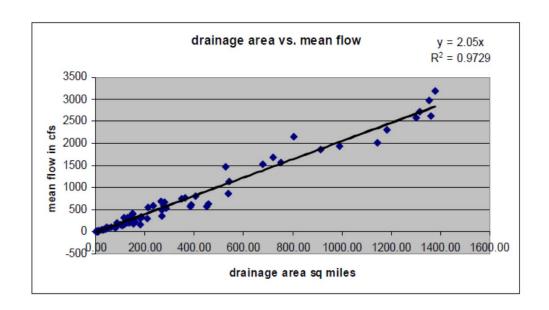
Q: event streamflow

Q_a Mean annual streamflow

Easily determined for gauged watersheds given sufficiently long records (> 20 yrs.);

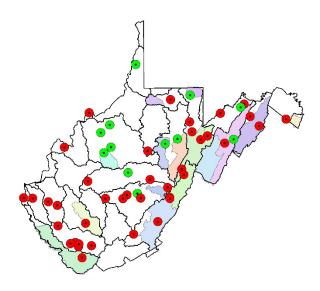
More difficult for ungauged watersheds;

Previous studies have used regression model f (drainage area);

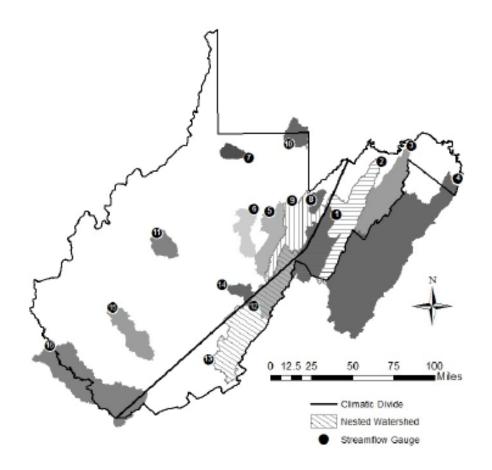




Big Sandy Creek USGS Streamflow station

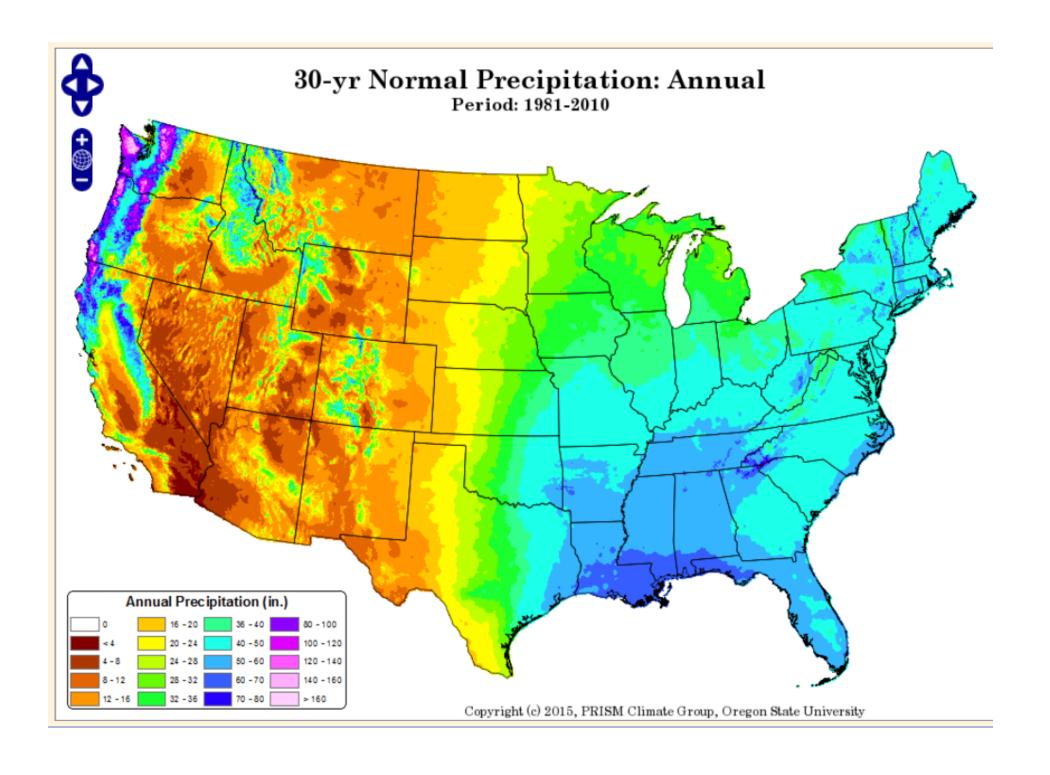


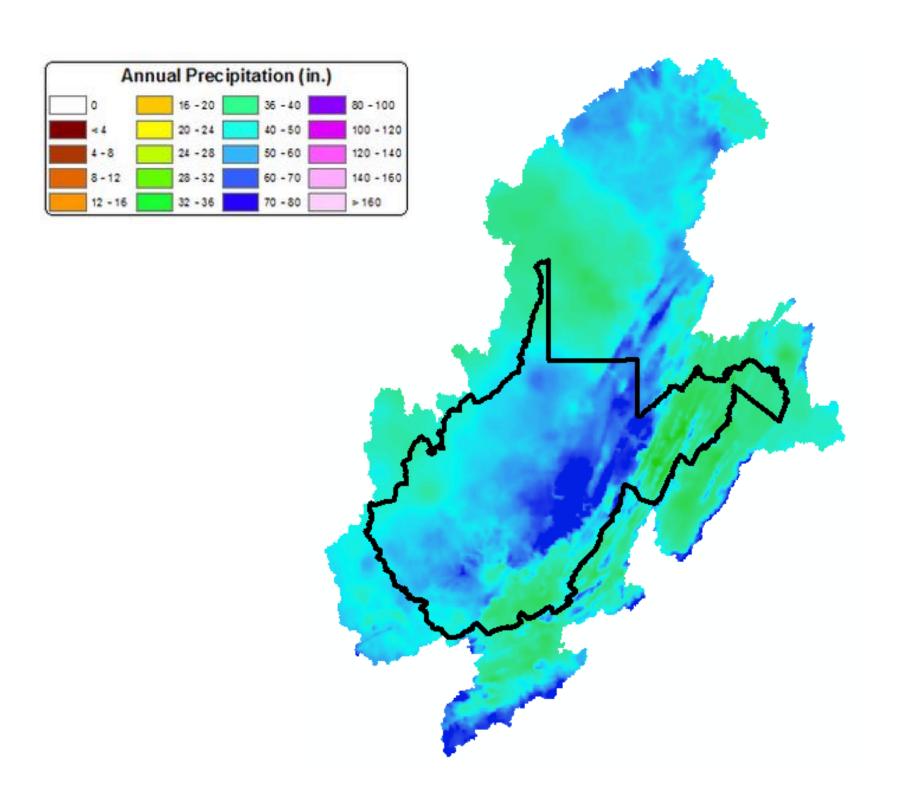
USGS streamflow stations in WV



1	South Branch Potomac River near Petersburg	9	Cheat River near Parsons
2	South Branch Potomac River near Springfield	10	Big Sandy Creek at Rockville
3	Cacapon River near Great Cacapon	11	West Fork Little Kanawha River at Rocksdale
4	Shenandoah River at Millville	12	Greenbrier River at Buckeye
5	Tygart Valley River at Belington	13	Greenbrier River at Alderson
6	Buckhannon River at Hall	14	Williams River at Dyer
7	Buffalo Creek at Barrackville	15	Big Coal River at Ashford
8	Blackwater River at Davis	16	Tug Fork at Kermit

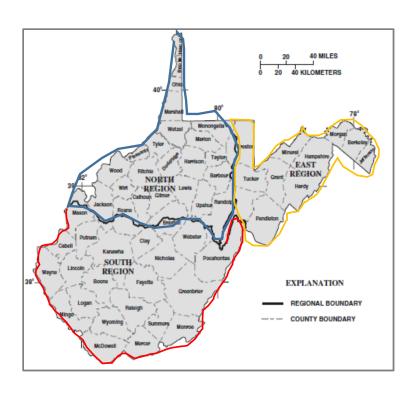
Figure 1. Locations of the 16 study watersheds throughout West Virginia.





Updated version uses Q/P relationship from regional modeled outputs

 Limited by the free flowing upstream contributing area above a USGS gauge site with records that match the precipitation data



Region	Watersheds	USGS # P-Q Regression Model	R^2	Q/P
North	Big Sandy	$3070500 \ y = 0.9691x - 503.69$	0.888	0.57
North	Blackwater	$3066000 \ y = 0.9512x - 496.25$	0.9	0.6
North	Buckhannon	$3053500 \ y = 0.9529x - 564.02$	0.92	0.55
North	Buff Ck	$3061500 \ y = 0.8272x - 507.36$	0.855	0.4
North	Cheat	3069500 y = 1.015x - 539.03	0.939	0.63
North	Tygart Val	$3051000 \ y = 0.9669x - 546.51$	0.929	0.55
North	William	$3186500 \ y = 0.8876x - 417.14$	0.83	0.61
South	GB Ald	3183500 y = 0.9696x - 546.61	0.888	0.46
South	GB Buck	3182500 y = 0.9459x - 566.26	0.888	0.48
South	Big Coal	$3198500 \ y = 0.7698x - 464.09$	0.84	0.38
South	TugKermit	3214500 y = 0.6861x - 393.82	0.787	0.28
East	SB Pot Peters	$1606500 \ y = 1.7037x - 992.39$	0.76	0.73
East	SB Pot Spring	$1608500 \ y = 0.8094x - 425.27$	0.797	0.35
East	ShenMillville	$1636500 \ y = 0.6684x - 344.43$	0.688	0.31

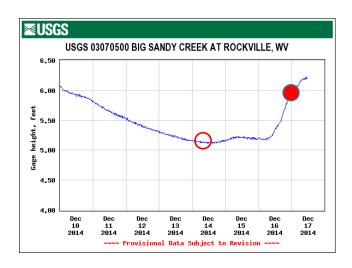
Determining Q event streamflow

$$Q_a' = \frac{Q}{Q_a}$$

Previous ZCC version used bankfull flow calculated as 90% of annual flow

Updated version uses annual peakflow streamflows from *Wiley et al., 2002*

 Used to represent a range of flow conditions from 1.1 to 3-yr recurrence storms.



U.S. Department of the Interior U.S. Geological Survey

Estimating the Magnitude of Annual Peak Discharges with Recurrence Intervals between 1.1 and 3.0 Years for Rural, Unregulated Streams in West Virginia

By JEFFREY B. WILEY, JOHN T. ATKINS, JR., and DAWN A. NEWELL
Water-Resources Investigations Report 02-4164

in cooperation with the WEST VIRGINIA DEPARTMENT OF TRANSPORTATION, DIVISION OF HIGHWAYS; WEST VIRGINIA SOIL CONSERVATION AGENCY; and the WEST VIRGINIA GEOLOGICAL AND ECONOMIC SURVEY

Charleston, West Virginia 2002

Determining Q event streamflow

$$Q'_a = \frac{Q}{Q_a}$$

Tirginia.

Table 1. Equations and regression statistics determined in the regional regression analysis of peak discharges

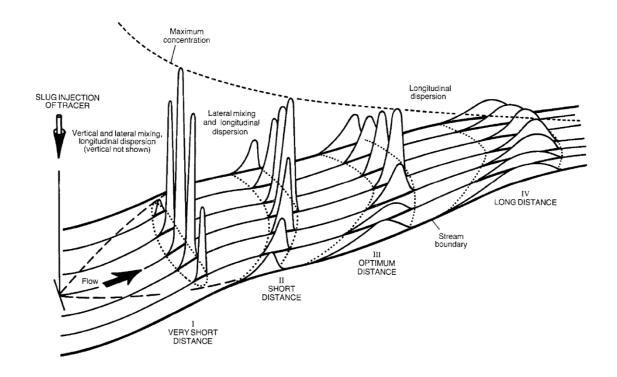
[Q(n) is the discharge in cubic feet per second for the (n)-year recurrence interval; A is the drainage area in square miles]

Estimating the Magnitude of Annual Peak Discharges with Recurrence Is between 1.1 and 3.0 Years for **Unregulated Streams in**

Regression equation	Standard error of the model, in percent	Average standard error of sampling, in percent	Average prediction error, in percent	Equivalent years of record	Number of streamflow stations	Range of	Tirginia LEY, JOHN T. ATKINS, JR., a estigations Report 02-4164	nd DAWN A. NEWELL	
			East Region				•		
Q(1.1) = 31.7 A ^{0.834} Q(1.2) = 37.9 A ^{0.8} Q(1.3) = 42.6 A ^{0.8}	51.1	10.1	52.4	1.3 North Region		1			
$Q(1.4) = 46.5 \text{ A}^{0.8}$ $Q(1.5) = 49.9 \text{ A}^{0.8}$	Q(1.1) = 57.7 A ^{0.789} Q(1.2) = 73.2 A ^{0.771}	30.8	7.6	31.8	2.6 So	uth Region		ı	
$Q(1.7) = 53.0 \text{ A}^{0.8}$ $Q(1.7) = 55.8 \text{ A}^{0.8}$ $Q(1.8) = 58.4 \text{ A}^{0.8}$ $Q(1.9) = 60.9 \text{ A}^{0.8}$ $Q(2) = 62.6 \text{ A}^{0.842}$	Q(1.3) = 85.3 A 0.759 Q(1.4) = 95.5 A 0.751 Q(1.5) = 105 A 0.744 Q(1.6) = 113 A 0.738 Q(1.7) = 120 A 0.733 Q(1.8) = 127 A 0.728	Q(1.1) = 46 Q(1.2) = 56 Q(1.3) = 64 Q(1.4) = 70 Q(1.5) = 75 O(1.6) = 80 Q(1.7) = 84 Q(1.8) = 88 Q(1.9) = 92 Q(2) = 95.4	.9 A 0.799 .4 A 0.795 .5 A 0.792 .8 A 0.790 .6 A 0.789 .8 A 0.787 .8 A 0.785	47.4 44.0 42.3 41.1 40.3 39.7 39.2 38.8 38.5 38.4	8.0 7.6 7.6 7.6 7.3 7.3 7.3 7.3 7.3 7.3	48.3 44.8 43.1 41.9 41.1 40.5 40.0 39.6 39.3 39.2	1.1 1.3 1.4 1.4 1.5 1.5 1.5 1.6 1.6	100	0.10-8,371
		Q(2.5) = 11 Q(3) = 121	0 A ^{0.781} A ^{0.778}	37.6 37.2	7.3 6.9	38.4 38.0	1.7 1.7		

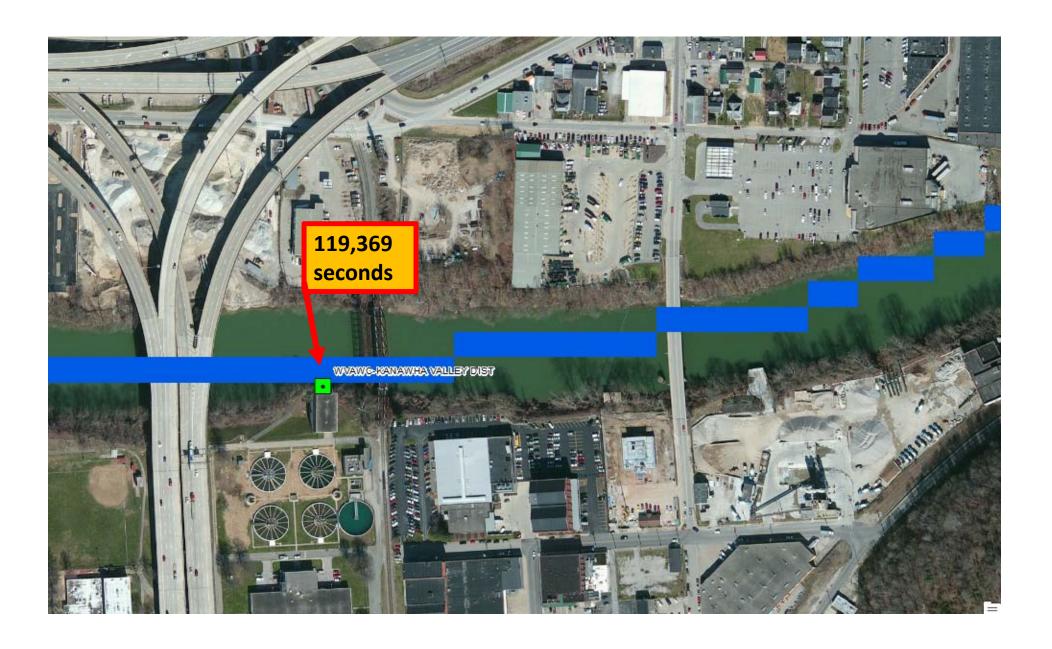
$$V_{mp} = 0.82 + \left(0.02 * D_a^{\prime 0.919} \ Q_a^{\prime -0.469} \ S^{0.159} \ \frac{Q}{D_a}\right)$$

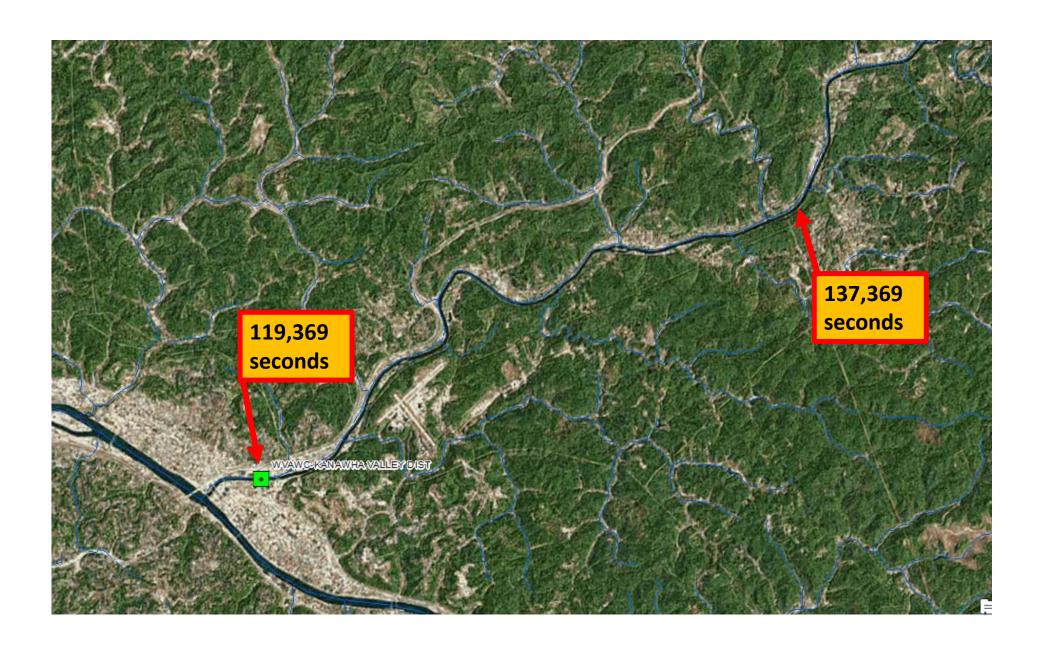
 V_{mp} = Maximum probable velocity

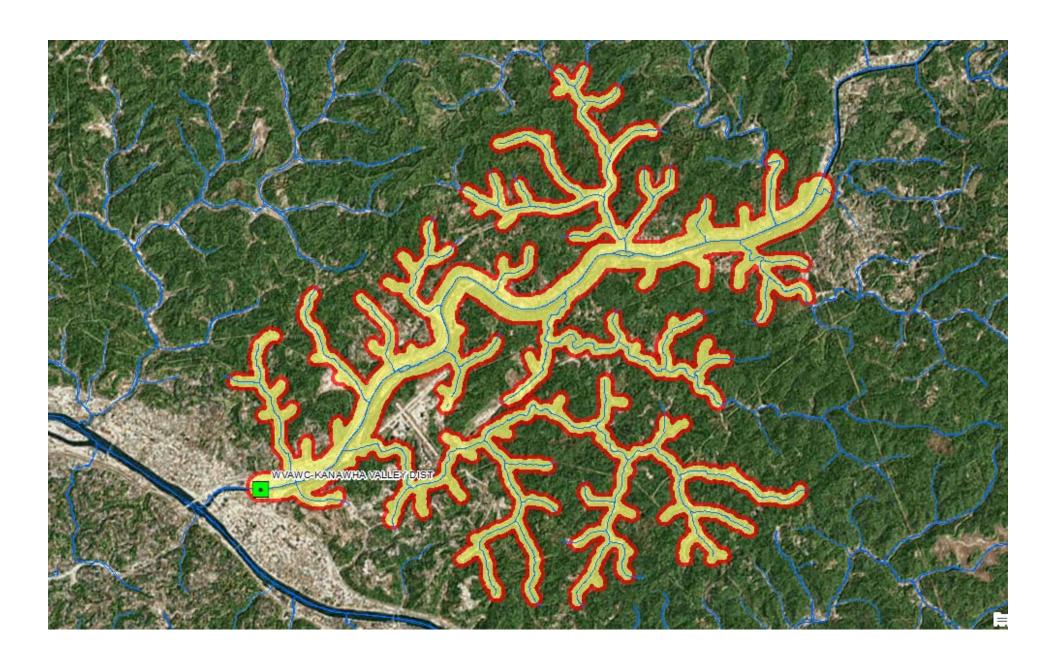


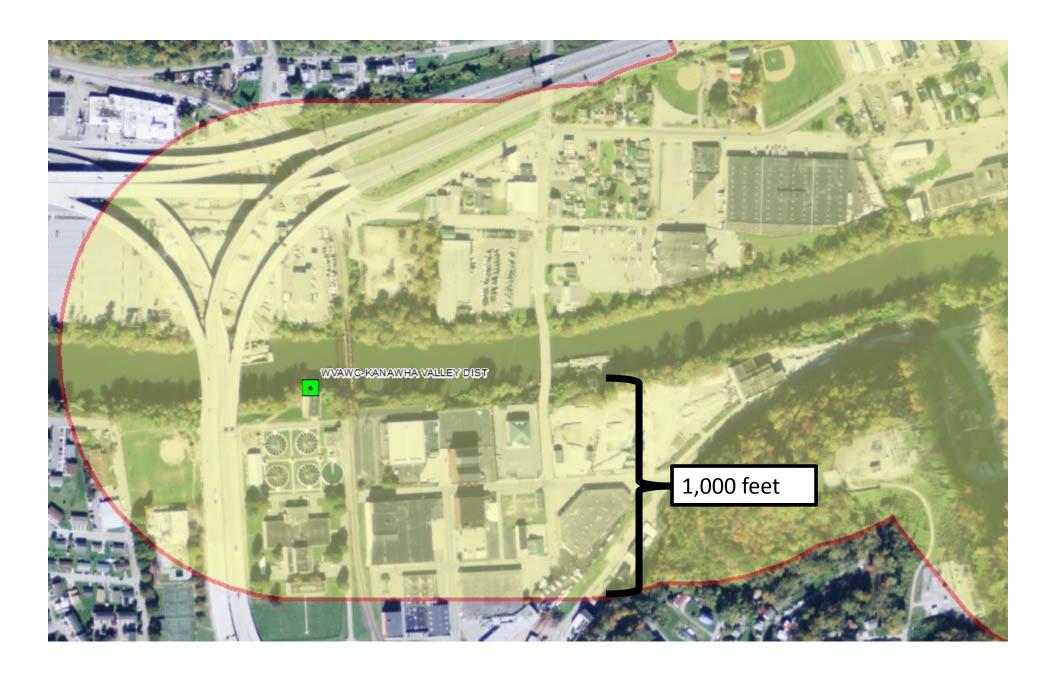


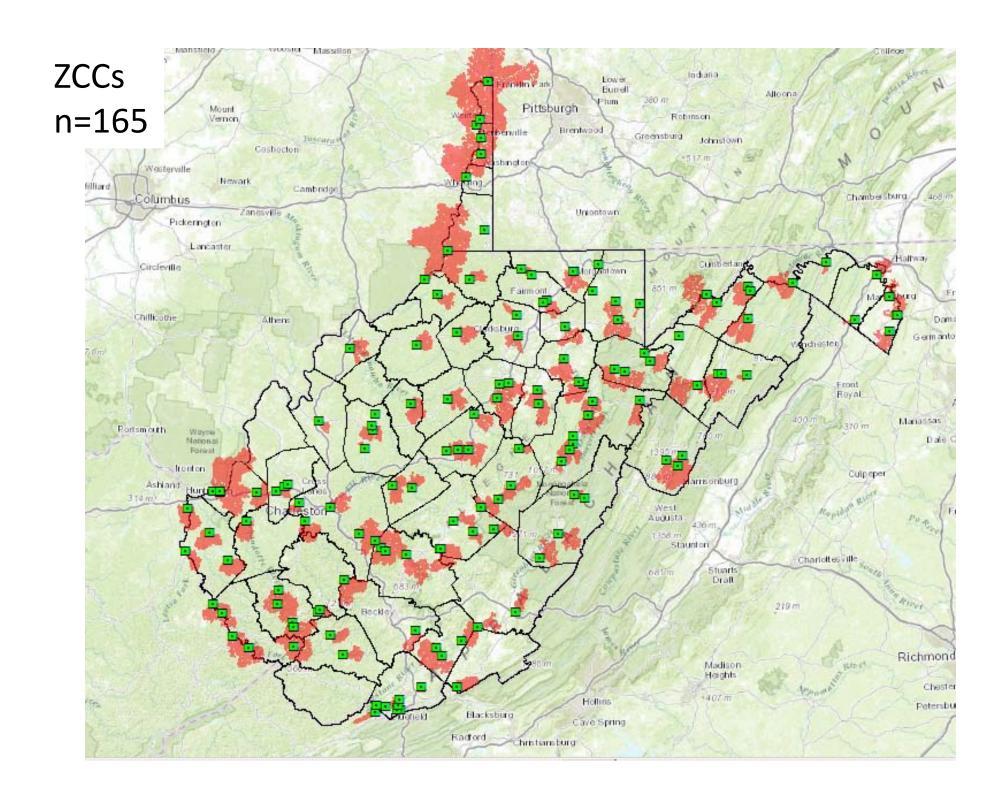


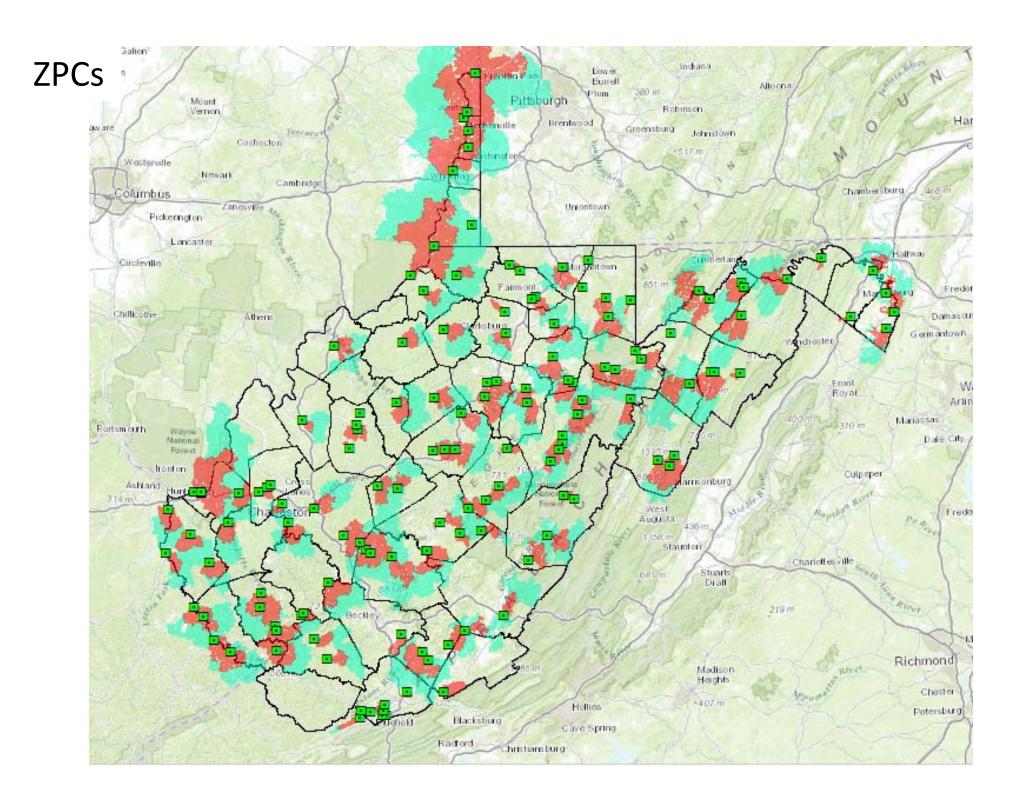












Limitations

- Three streamflow regions
- Doesn't consider dispersion plume for specific chemicals
- Maximum velocity equation is very conservative
- Lack of calibration or validation at this time





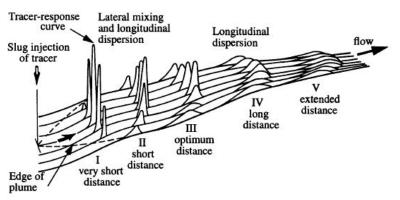


Figure 1. Lateral mixing and longitudinal dispersion patterns and changes in distribution of concentration downstream from a single, center, slug injection of tracer. (Modified from Kilpatrick, 1993, p. 2.)

Why the 2001 and 2015 ZCC differences?

- 100K streams versus 24k streams
- One stream flow equation versus three regional ones
- Bankfull flow of 90% of high flow versus statistical occurrence of 90% flow
- 30m elevation data versus 1 and 3m elevation data to provide drainage area and stream gradient
- Buffers from stream centerline versus stream banks